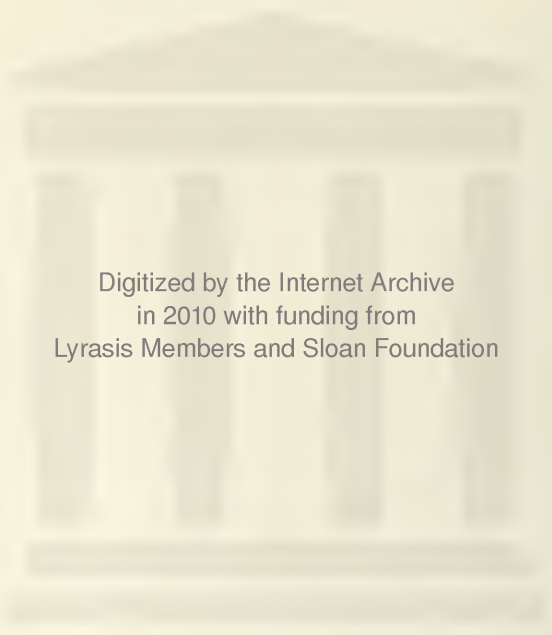


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NORTHEAST REGIONAL RESEARCH PUBLICATION

Drought Frequency in the Northeastern United States

BULLETIN 595

WEST VIRGINIA UNIVERSITY

JUNE 1970

AGRICULTURAL EXPERIMENT STATION



The Authors

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Technical Committee of Northeastern Regional Research
Project NE-35, "Climate of the Northeast — Analysis and
Relationship to Plant Response"

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Summary

Studies by Fieldhouse and Palmer (4) have established that the Northeast United States, a humid region, has been subjected to numerous droughts in the past 40 years. The Drought Index developed by Palmer (8) provides a measure of drought severity and duration. Employing this concept the frequency, expressed as a return period or recurrence interval, for drought severity was determined for the Northeast.

In terms of frequency, the general expectancy is for moderate drought (-2.00 to -2.99) once in five years, severe (-3.00 to -3.99) once in ten years, and a severity of -5.00 to -7.00 once in 50 years. The stated values of the Index are expected on the average to be equalled or exceeded once in the periods given if no periodicity is implied. While these expectancies generally apply to the Northeast there is significant variation within the region.

Areas of least drought severity include the mountainous parts of southwestern Virginia and the western Carolinas, some coastal areas of the Carolinas, part of New York adjacent to the Great Lakes and all of New England east of the Connecticut Valley. Areas of worst severity include the Hudson Valley, the Del-Mar-Va Peninsula, eastern Maryland, northern Virginia, northern West Virginia along with some climatic divisions in southwestern Pennsylvania and eastern Ohio. With the exception of the coincidence of low severity with the high rainfall areas of the mountains of Virginia and the Carolinas there is a striking relationship between drought patterns and topography or meteorological characteristics of the Northeast.

Drought Frequency in the Northeastern United States

W. H. DICKERSON AND B. E. DETHIER

THE VARIABILITY of precipitation is a characteristic of the climate of the Northeast as well as most other humid areas. It is not unusual for areas of seemingly abundant precipitation to suffer a soil moisture shortage that affects agricultural production and the deficit of precipitation not infrequently persists to the point where groundwater levels and stream flow are adversely influenced. To describe these situations numerous definitions of drought have been proposed but none has received general or widespread acceptance. A few examples of the many that may be found in the literature, Hoyt (6), Havens (5), Carr (1) and Fieldhouse and Palmer (4), serve to illustrate the lack of agreement and, in some cases subjectivity, involved in establishing the existence of drought. It is nevertheless an inescapable fact that drought is a troublesome and costly feature of the climate of the Northeastern region of the United States as well as other humid regions of the world. Reference (9) contains papers on aspects of the drought of the 1960's in the Northeast.

Effects on Agriculture, Municipalities and Industry

An agricultural drought is experienced when rainfall is inadequate to maintain soil moisture at optimum levels in terms of the normal or generally experienced moisture supply for the region in question. Thus the meaning of adequate soil moisture is relative. However, crop production is related to plant species, soil fertility, and other environmental factors in addition to soil moisture. Lack of available soil moisture for crop production is generally the first manifestation of a precipitation deficit, therefore agriculture is particularly sensitive to drought. Indeed, a period of only 10 to 14 days without rain may drastically curtail germination of seeds or the growth or maturation of some crops on some soils. The more shallow rooted the crop and the smaller the available water holding capacity of the soil the greater is the vulnerability to drought.

A meteorological drought may be described, e.g., Carr (1), as a significant decrease from normal precipitation over a wide area and for an extended time. This situation is more severe and widespread than the local or short-term variation in rainfall that can create an agricultural drought. Hydrologic drought

is still more widespread and severe, usually lasting one or more years. It is typified by the drying up of springs and small streams, falling water levels in wells, the shrinking of rivers, and the depletion of water stored in lakes and reservoirs. Yevjevich (9) has proposed criteria for objective definitions of hydrologic droughts.

A drought may exist for the agriculturalist before it is evident in the meteorological or hydrological sense. Conversely an agricultural drought may be ended, at least temporarily, by rainfall that replenishes the available soil moisture supply but does not add to groundwater or appear as stream flow. An agricultural drought may exist sporadically because of poor distribution of rainfall when the total precipitation for the year is near normal or above and no meteorological or hydrological drought would be evident. If an agricultural drought may exist without serious consequences for municipal and industrial interests, it is likewise true that the latter may run low on water when agriculture does not suffer. This can happen when rainfall is so fortuitously distributed as to keep the available soil moisture for crops in a reasonable balance with the demands of evapotranspiration but such precipitation episodes may provide little or no excess moisture for the replenishment of groundwater or contribution to surface runoff. Agriculture needs water mostly during the warm or growing season. In contrast to this, municipal and industrial water supplies undergo withdrawal on a year-around basis with little seasonal diminution in requirements.

As these definitions and concepts imply, drought is harmful and costly to many segments of the economy but the most directly affected are agriculture and municipal and industrial users of water. Agriculturists are primarily concerned because of the importance of soil moisture in crop production. Engineers and hydrologists are interested in drought because of its effect upon water quality, streamflow, waste removal, and ground water levels.

Causes of Drought

Many theories have been advanced to explain the onset and persistence of drought. Some of the current ideas received with the most credence suggest that changes in the rainfall regime may be due to colder than usual temperatures of the oceans around North America, changes in the large-scale circulation of air resulting in a shift of prevailing wind patterns, low sunspot activity, changes in the climate caused by dust thrown into the atmosphere by volcanic eruptions. An underlying concept common to most theories is that variations in the energy from the sun are directly or indirectly responsible for changes in the climate.

Synoptic patterns accompanying the drought of the 1960's in the Northeast indicate a slowdown or failure in the flow of moisture laden air from the Gulf of Mexico or the Atlantic Ocean, a change in the movement of frontal systems across the region, or the failure of coastal storms to follow a tract that brings moisture into the area. Also, there was a greater frequency of anticyclones moving across the Northeastern United States during the warm months. The more frequent invasions of cool dry air interfered with the normal flow of moist tropical air from the Gulf of Mexico.

Predicting Drought

Due to the complexity of the atmospheric circulation long-range forecasts have lacked the accuracy necessary for drought prediction. As a result, interest has long centered about the identification of cycles in the weather that could be a clue to the expected timing of occurrence, severity and length of drought periods. Indeed, this aspect of the weather has been so fascinating that some 130 cycles, Tannehill (10), have been proposed at one time or another for explaining the vagaries of the weather. One of the most well-known of these cycles is the Luckner, attributable to the Viennese climatologist who proposed a 35-year cycle. Another that has received much attention is the sunspot cycle which averages approximately 11 years but has been as short as 7 or as long as 17 years. Somewhat along this same line, Tannehill explained how current weather sometimes seems to duplicate previous weather development and cited examples of how analogs have been used in forecasting. Of the many cycles which have been proposed, none offers a practical means of forecasting the beginning, severity or end of a drought period.

Weather Modification

In relation to drought, interest in weather modification centers around two divergent concerns, namely, inadvertent modification, and a controlled or scientifically induced modification for beneficial purposes. As an example of the first concern, it has been suggested that air pollution may be a factor tending to bring about unanticipated and uncontrolled changes in the climate. These may include a shift in rainfall patterns, but the present understanding of the variables involved does not provide the basis for a satisfactory explanation or prediction of the changes to be expected. In the past two decades, the interest in "weather making" has spread throughout the United States and indeed all over the world. Based on ideas advanced by Drs. Langmuir, Schaefer, Vonnegut, and others, developments in the late 1940's raised some hope that weather modification by cloud seeding could bring widespread relief from drought, increase precipitation over arid areas, eliminate lightning and hail as destructive

manifestations of local weather and hopefully modify or steer tropical hurricanes. A large number of carefully planned studies are now in progress on various aspects of weather modifications (11). The results to date suggest that relief from the vagaries of the weather through this means is not imminent although holding out some hope of at least limited success for the future. The conclusion seems plain that protection from drought and augmentation of water supplies will have to be dependent on conventional methods for the immediate future.

Purpose of the Study

The work by Fieldhouse and Palmer (4) cited previously, indicates that the Northeast has always been subject to wet and dry spells. Commonly accepted means of alleviating the effects of drought include irrigation for agriculture and the storage of surface water in reservoirs and the drilling of more and deeper wells to tap groundwater for use by industry and for municipal consumption. These methods are expensive—so costly that the most careful scrutiny should be given to the various aspects of drought and the impact of drought on the water supply. As the demand for water increases, problems of evaluation, selection, and development of supplies will necessarily become more complex.

Whenever man makes long-range plans that depend on the weather, he must depend on the climatological and hydrological data for the area. Present knowledge allows accurate weather forecasts for periods up to about five days, but this is generally too short to be of much use in decisions involving drought. A knowledge of drought occurrence in a region over past years will allow estimation of the chances of occurrence in the future, a prerequisite in planning and designing for water supplies.

The Palmer Drought Index (8), which provides a measure of drought severity as well as duration, offered a unique opportunity to study drought frequency. The central purpose of the work reported herein has been to determine the frequency or return period associated with a given severity of drought. The return period is the average number of years within which a given event will be equalled or exceeded. No periodicity is implied in this use of the term.

The Data Sample

The data were obtained from the method of drought analysis developed by Palmer (8) and applied to the Northeastern United States by Fieldhouse and Palmer (4). The procedure derived monthly index values which are a measure of drought severity and identified the months affected. According to Fieldhouse and Palmer the procedure "treats drought severity as a function of accumulated

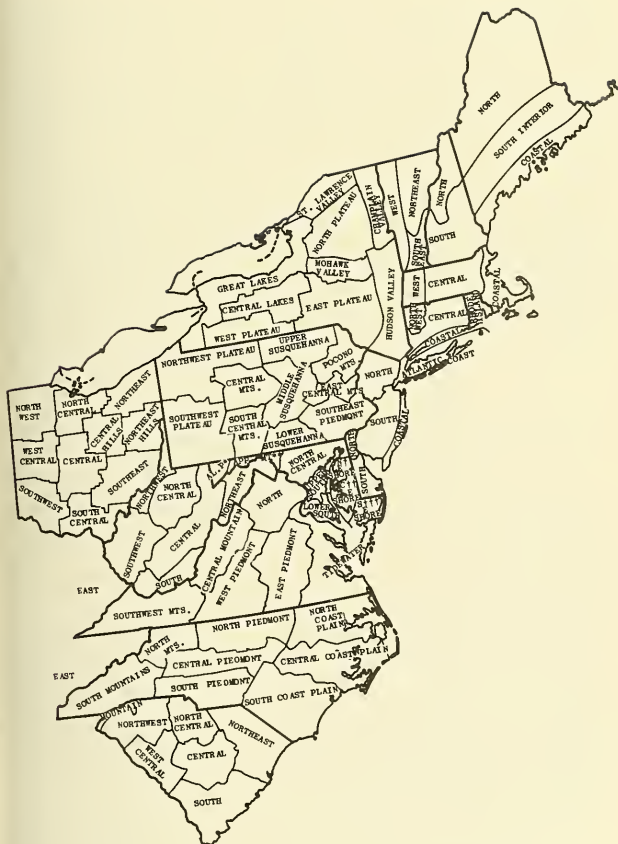


Figure 1. Climatological divisions used in the Northeast regional drought analysis.

weighted differences between actual precipitation and the precipitation requirement, where the requirement depends on the carryover of previous rainfall as well as on the evapotranspiration, moisture recharge and runoff that would be climatically appropriate for the particular time and place being investigated. Thus, the average requirement is for normal rainfall, but individual periods may require much above or much below normal rainfall depending on the character of the preceding weather and the temperature of the period in question."

The method did not attempt to consider biological responses to drought but concentrated on identification of various durations and severity of drought periods by an analysis of the historical weather records. These records were combined by climatic divisions so that the derived values apply to the geographical units delineated on the map of climatological divisions, (Figure 1). Categories of drought severity are defined in Table 1.

Table 1. Classes of Dry Periods According to Palmer (8).

| Drought Index | Description of Class |
|----------------|----------------------|
| -0.50 to -0.99 | incipient drought |
| -1.00 to -1.99 | mild drought |
| -2.00 to -2.99 | moderate drought |
| -3.00 to -3.99 | severe drought |
| ≤ -4.00 | extreme drought |

The Palmer Drought Index assigned positive values to wet periods and negative values to dry periods.

Analytical Procedures

Each drought period was identified by date of beginning and ending, the greatest severity index reached and by the duration in terms of consecutive months affected. Severity data, representing climatic divisions, were compiled into samples with the drought periods in chronological order. The data were arrayed in ascending order of severity with the data identifications for the periods retained. Data samples were then analyzed to obtain frequency relationships.

In this procedure, a drought was identified as existing when a specified index (≤ -0.50) was reached. Some years experienced no drought, others had one or more. For this reason the data samples were considered to constitute a partial series. The purpose of the study required establishment of a relationship between the event (severity index) and a frequency or return period. The return period was considered as the independent variable and the index the dependent

fitting positions were determined by the Weibull formula, Chow (2), which

$$F = \frac{M}{N+1}$$

F = the cumulative frequency which may also be expressed as a return period in years,

M = the order number of the event in an array,

N = the number of years in the record.

For the purpose of analyzing the partial series the order number was determined by assigning the lowest index (most extreme severity) order number of 1, i.e., $M = 1$. This gives the event a return period of $N + 1$ years and corresponds to the method of handling an annual series when the Weibull fitting equation is used.

The fitting procedure employed an exponential function in the form

$$y \ln b = \ln X - \ln a$$

This equation was reduced to a least squares fitting problem, using y as the index (severity) on a linear scale and X the frequency (return period) on a logarithmic scale (see Figure 2).

By graphical plots of the data and comparison of R^2 (coefficient of determination) values to determine the percentage of the variability in the event accounted for by the return period it was evident that this function generally provided a satisfactory fit to the data, with some exceptions noted later. The exponential function is supported as applying to partial series data by Linsley, Kohler and Paulhus (7) and by Chow (2). This model was subsequently used to determine the frequency relationships presented in Table 2.

Results of Frequency Analysis

The Region Studied

Drought severity data were derived by climatic divisions for the area from Maine to South Carolina and west to Ohio. Two climatic divisions were not included, North Carolina, North Mountains (31-02) and South Carolina Mountain (38-01) because they were not reported in the calculations by Fiedhouse and Palmer (4).

The chief distinguishing physical feature of the area is the Appalachian Mountains, running from Maine through the Carolinas. These mountains are bordered on the east by lands which slope gently down to the sea and on the west by more broken lands which slope away to lower elevations to the westward. The general precipitation pattern for the region has been described by

Dethier (3) as follows; "Most of the moisture for the region is transported from the Gulf of Mexico and the Atlantic Ocean by the major storm systems in the atmosphere. Although these storms are the major year round producers of precipitation there are fewer during the summer season. Increased convective activity over most of the area is more than adequate in compensating for the decrease in major storms. This results in a warm season precipitation maximum over the greater part of the region. The complex precipitation pattern is due, in part, to the rugged topography of the areas. Orographic lifting results in numerous centers of heavy precipitation over higher elevations and on the windward slopes. Areas of least precipitation are usually found in the lee or 'rain shadow' of the higher elevations.

The average annual precipitation ranges from 31 inches in extreme northwestern New York to more than 52 inches at some of the higher elevations in West Virginia, New York, Vermont and New Hampshire. . . The entire Atlantic seaboard area receives more than 44 inches per year as do sections of the Adirondacks, central West Virginia, and northwestern Pennsylvania. The areas with lowest values include the rain shadows in Maryland, New York, Pennsylvania and West Virginia."

Drought Severity

Tabulated values of the severity index for selected return periods are shown in Table 2, which also lists the greatest severity encountered for the period analyzed (1929-1967). For all of the climatic divisions the one in two year index value falls in the mild drought class (-1.00 to -1.99). With few exceptions the five-year expectancy is for a moderate drought category (-2.00 to -2.99) to be equalled or exceeded. The ten-year index shows that severe drought (-3.00 to -3.99) can be expected to occur within the span of a decade, on the average, at most locations in the northeastern region as defined in this study. Generally the 50-year values of the index are between -5.00 and -7.00, and it may be noted that these agree fairly closely with the extreme events observed although there are a few notable exceptions. For the most part the fitted line for the model agrees well with the observed data. Figure 2 illustrates the relationship for the Central division (19-02) of Massachusetts.

Regional Patterns

Table 2 gives values of calculated severity indices for the Northeast climatic divisions, while Figure 3 was designed to show the enveloping curves for the entire region. A question that arises naturally is, Do the climatic division values reported delineate a geographical pattern that may be related to topography, seasonal or annual rainfall, or other identifiable meteorological characteristics of the region? In examining this the calculated return period

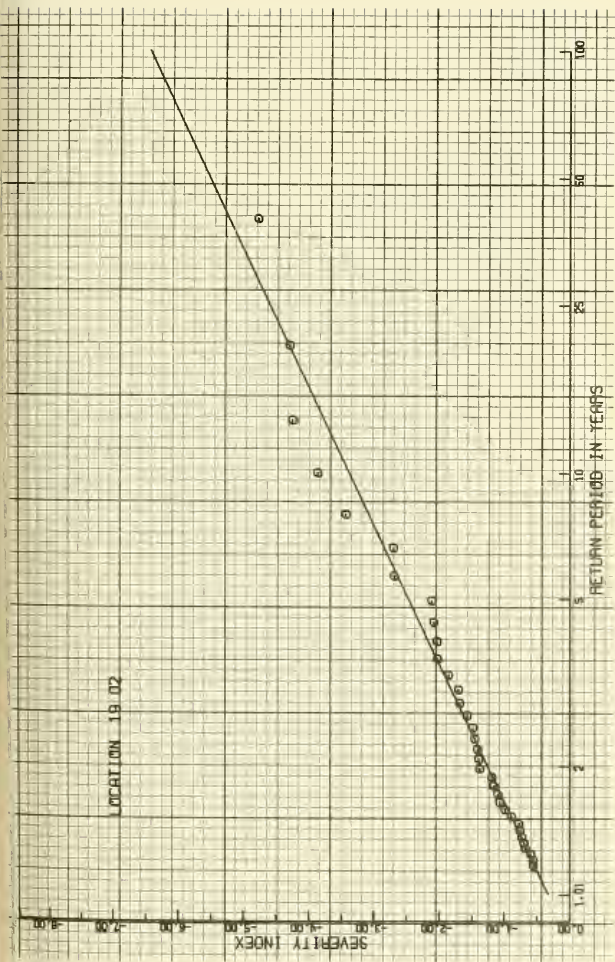
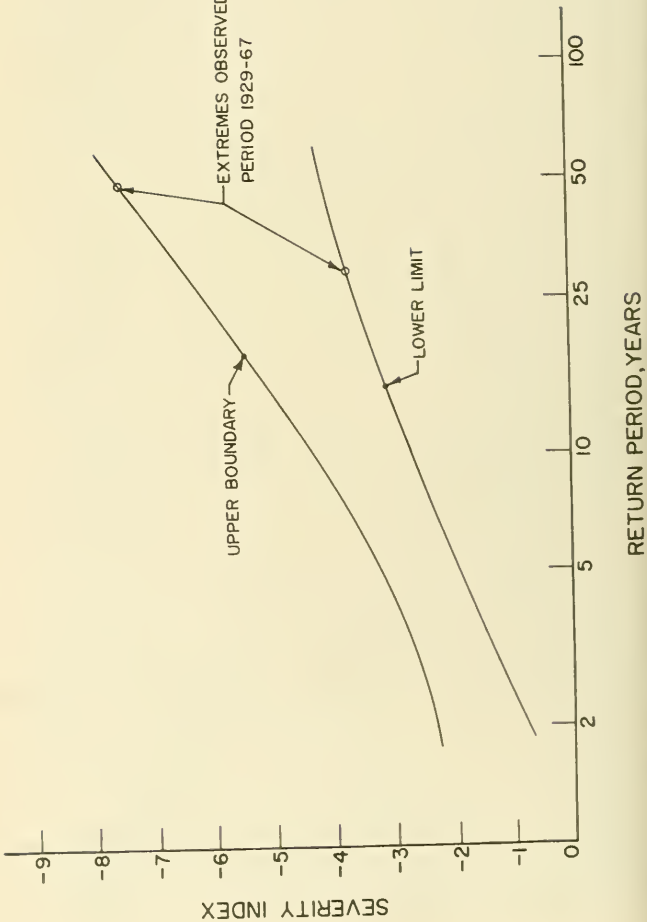


Figure 2. Frequency relationships for drought severity-Central climatological division of Massachusetts (19-02).



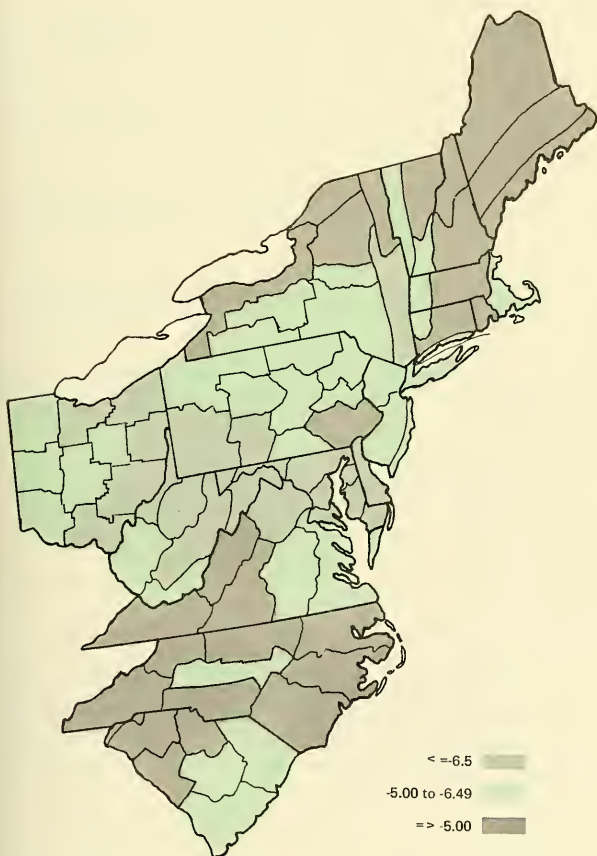


Figure 4. Greatest drought severity for period 1929-67.

TABLE 2. DROUGHT SEVERITY INDICES*

| State | Location | Division | Return Period Values | | | | Extreme observed | |
|--------------------|----------|----------------------------|----------------------|-------|--------|--------|---------------------|--------|
| | | | 2 Yr. | 5 Yr. | 10 Yr. | 25 Yr. | | 50 Yr. |
| Connecticut (06) | | Northwest (01) | -1.05 | -2.37 | -3.37 | -4.67 | -5.68 | -5.44 |
| | | Central (02) | -1.30 | -2.43 | -3.28 | -4.41 | -5.26 | -4.61 |
| | | Coastal (03) | -1.36 | -2.62 | -3.58 | -4.84 | -5.80 | -5.67 |
| Delaware (07) | | North (01) | -1.04 | -2.77 | -4.07 | -5.80 | -7.10 | -6.54 |
| | | South (02) | -0.87 | -2.72 | -4.11 | -5.95 | -7.35 | -6.66 |
| Maine (17) | | North (01) | -1.31 | -2.37 | -3.17 | -4.22 | -5.02 | -4.49 |
| | | South Interior (02) | -1.23 | -2.35 | -3.20 | -4.32 | -5.17 | -4.27 |
| | | Coastal (03) | -1.34 | -2.58 | -3.52 | -4.77 | -5.72 | -4.95 |
| Maryland (18) | | South Eastern Shore (01) | -1.06 | -2.66 | -3.88 | -5.48 | -6.69 | -7.02 |
| | | Central Eastern Shore (02) | -1.17 | -2.81 | -4.04 | -5.67 | -6.91 | -6.90 |
| | | Lower South (03) | -1.30 | -2.91 | -4.13 | -5.74 | -6.97 | -7.03 |
| | | Upper South (04) | -1.21 | -2.88 | -4.13 | -5.79 | -7.05 | -7.22 |
| | | North Eastern Shore (05) | -1.01 | -2.86 | -4.26 | -6.11 | -7.51 | -6.75 |
| | | North Central (06) | -1.24 | -2.93 | -4.20 | -5.89 | -7.16 | -7.10 |
| | | Appalachian Mts. (07) | -1.36 | -2.93 | -4.12 | -5.70 | -6.89 | -6.44 |
| | | Allegheny Plateau (08) | -1.43 | -2.81 | -3.85 | -5.23 | -6.27 | -5.84 |
| | | West (01) | -1.04 | -2.67 | -3.91 | -5.53 | -6.77 | -6.53 |
| Massachusetts (19) | | Central (02) | -1.23 | -2.48 | -3.42 | -4.67 | -5.61 | -4.78 |
| | | Coastal (03) | -1.23 | -2.48 | -3.42 | -4.67 | -5.61 | -5.09 |
| | | North (01) | -1.12 | -2.30 | -3.20 | -4.38 | -5.27 | -4.23 |
| New Hampshire (27) | | South (02) | -1.32 | -2.49 | -3.38 | -4.55 | -5.44 | -4.66 |
| | | North (01) | -1.24 | -2.49 | -3.43 | -4.68 | -5.62 | -5.32 |
| New Jersey (28) | | South (02) | -1.42 | -2.80 | -3.85 | -5.23 | -6.28 | -6.47 |
| | | Coastal (03) | -1.40 | -2.62 | -3.53 | -4.76 | -5.68 | -5.09 |

New York (30)

| | | | | | | |
|----------------------------|--------------|-------|-------|-------|-------|-------|
| West Plateau (01) | -1.22 | -2.70 | -3.82 | -5.30 | -6.42 | -5.39 |
| East Plateau (02) | -1.28 | -2.65 | -3.68 | -5.06 | -6.09 | -5.85 |
| North Plateau (03) | -1.32 | -2.50 | -3.40 | -5.48 | -5.47 | -4.13 |
| Coastal (04) | -1.36 | -2.64 | -3.60 | -4.88 | -5.85 | -6.22 |
| Hudson Valley (05) | -1.29 | -2.76 | -3.87 | -5.34 | -6.45 | -6.81 |
| Mohawk Valley (06) | -1.56 | -2.76 | -3.67 | -4.87 | -5.79 | -5.21 |
| Champlain Valley (07) | -1.45 | -2.82 | -3.85 | -5.21 | -6.24 | -4.62 |
| St. Lawrence Valley (08) | -1.28 | -2.56 | -3.52 | -4.80 | -5.76 | -4.52 |
| Great Lakes (09) | -1.34 | -2.49 | -3.36 | -4.51 | -5.38 | -4.51 |
| Central Lakes (10) | -1.19 | -2.57 | -3.62 | -5.00 | -6.05 | -5.03 |
| South Mountains (01) | -1.38 | -2.17 | -2.67 | -3.54 | -4.13 | -3.72 |
| North Mountains (02) | Not reported | | | | | |
| North Piedmont (03) | -1.36 | -2.48 | -3.33 | -4.45 | -5.30 | -4.76 |
| Central Piedmont (04) | -1.29 | -2.40 | -3.24 | -4.35 | -5.20 | -5.47 |
| South Piedmont (05) | -1.33 | -2.38 | -3.18 | -4.23 | -5.03 | -4.96 |
| South Coastal Plain (06) | -0.94 | -2.07 | -2.92 | -4.05 | -4.90 | -4.06 |
| Central Coastal Plain (07) | -1.13 | -2.38 | -3.33 | -4.59 | -5.54 | -4.04 |
| North Coastal Plain (08) | -1.26 | -2.26 | -3.02 | -4.02 | -4.78 | -4.53 |
| Northwest (01) | -0.75 | -2.70 | -4.17 | -6.11 | -7.59 | -5.87 |
| North Central (02) | -0.74 | -2.71 | -4.20 | -6.18 | -7.67 | -6.61 |
| Northeast (03) | -0.84 | -2.39 | -3.56 | -5.10 | -6.27 | -6.57 |
| West Central (04) | -1.08 | -2.76 | -4.02 | -5.69 | -6.95 | -6.10 |
| Central (05) | -1.09 | -2.91 | -4.28 | -6.09 | -7.46 | -6.21 |
| Central Hills (06) | -1.42 | -3.08 | -4.34 | -5.99 | -7.25 | -6.12 |
| Northeast Hills (07) | -0.98 | -2.73 | -4.05 | -5.81 | -7.14 | -6.60 |
| Southwest (08) | -1.29 | -3.02 | -4.32 | -6.05 | -7.36 | -6.03 |
| South Central (09) | -1.16 | -2.91 | -4.23 | -5.98 | -7.31 | -6.81 |
| Southeast (10) | -1.43 | -2.93 | -4.06 | -5.57 | -6.70 | -7.25 |

North Carolina (31)

Ohio (33)

Table 2 con't

| | | | | | | | |
|---------------------|------------------------------|--------------|-------|-------|-------|-------|-------|
| Pennsylvania (33) | Pocomo Mts. (01) | -1.16 | -2.52 | -3.54 | -4.90 | -5.92 | -5.70 |
| | East Central Mts. (02) | -1.38 | -2.69 | -3.68 | -4.99 | -5.97 | -5.23 |
| | Southeast Piedmont (03) | -1.20 | -2.57 | -3.61 | -4.98 | -6.01 | -4.96 |
| | Lower Susquehanna (04) | -1.29 | -2.82 | -3.98 | -5.52 | -6.68 | -6.21 |
| | Middle Susquehanna (05) | -1.05 | -2.55 | -3.68 | -5.17 | -6.30 | -6.42 |
| | Upper Susquehanna (06) | -1.04 | -2.48 | -3.56 | -4.99 | -6.08 | -5.69 |
| | Central Mountains (07) | -1.34 | -2.59 | -3.53 | -4.78 | -5.72 | -5.78 |
| | South Central Mountains (08) | -1.55 | -3.01 | -4.12 | -5.58 | -6.69 | -7.18 |
| | Southwest Plateau (09) | -1.42 | -2.83 | -3.89 | -5.29 | -6.35 | -6.70 |
| | Northwest Plateau (10) | -1.15 | -2.31 | -3.19 | -4.35 | -5.22 | -5.93 |
| Rhode Island (37) | | -1.13 | -2.36 | -3.29 | -4.51 | -5.44 | -4.67 |
| South Carolina (38) | Mountain (01) | not reported | | | | | |
| | Northwest (02) | -1.14 | -2.18 | -2.97 | -4.01 | -4.79 | -4.58 |
| | North Central (03) | -1.21 | -2.28 | -3.09 | -4.16 | -4.97 | -4.69 |
| | Northeast (04) | -0.97 | -2.25 | -3.22 | -4.51 | -5.48 | -5.08 |
| | West Central (05) | -1.20 | -2.24 | -3.02 | -4.06 | -4.85 | -3.93 |
| | Central (06) | -0.92 | -2.14 | -3.07 | -4.29 | -5.21 | -5.00 |
| | Southern (07) | -0.75 | -2.10 | -3.16 | -4.53 | -5.57 | -5.41 |
| | Northeast (01) | -1.31 | -2.42 | -3.26 | -4.37 | -5.21 | -4.21 |
| | West (02) | -1.15 | -2.54 | -3.60 | -4.99 | -6.04 | -5.69 |
| | Southeast (03) | -1.11 | -2.39 | -3.35 | -4.63 | -5.60 | -5.72 |
| Virginia (44) | Tidewater (01) | -1.02 | -2.57 | -3.75 | -5.30 | -6.48 | -5.94 |
| | East Piedmont (02) | -1.06 | -2.46 | -3.53 | -4.93 | -6.00 | -5.13 |
| | West Piedmont (03) | -1.35 | -2.56 | -3.47 | -4.68 | -5.59 | -4.49 |
| | North (04) | -1.25 | -3.90 | -4.15 | -5.80 | -7.05 | -7.39 |
| | Central Mts. (05) | -1.26 | -2.47 | -3.38 | -4.59 | -5.51 | -4.49 |
| | Southwest Mts. (06) | -1.27 | -2.65 | -3.68 | -5.06 | -6.09 | -4.90 |
| | Northwest (01) | -1.16 | -2.90 | -4.21 | -5.95 | -7.26 | -7.14 |
| | North Central (02) | -1.37 | -2.99 | -4.22 | -5.84 | -7.07 | -6.96 |
| | Southwest (03) | -1.05 | -2.78 | -4.08 | -5.80 | -7.10 | -6.22 |
| | Central (04) | -1.43 | -3.00 | -4.18 | -5.75 | -6.93 | -6.52 |
| West Virginia (46) | South (05) | -1.29 | -2.77 | -3.90 | -5.38 | -6.51 | -6.17 |
| | Northeast (06) | -1.32 | -3.02 | -4.31 | -6.02 | -7.30 | -7.46 |

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values of the index were not used. Instead the most extreme drought conditions countered were plotted by climatic divisions and are shown in Figure 4.

Areas of least drought severity, as measured by the Palmer Index, include the mountainous portion of southwestern Virginia and the Carolinas, some coastal areas of the Carolinas, the area in New York adjacent to the lakes and generally all of New England which lies to the east of the Connecticut River. Areas which have experienced the worst severity included parts of the Hudson Valley, the Del-Mar-Va peninsula, eastern Maryland, northern Virginia, northern West Virginia, along with the extreme southwestern divisions of Pennsylvania and the eastern parts of Ohio. Intermediate areas for severity included some tidewater areas of the Carolinas and Virginia, upper Pennsylvania and lower New York and this category also encompassed New Jersey and the coastal division of New York. As might be expected this pattern showed some resemblance to the normal annual rainfall but the correspondence is not very striking except that the least severe droughts were associated with the highest elevations and high rainfall areas of the Southern Appalachians.

With the exception of the coincidence of low severity with the high rainfall areas of the mountains of Virginia and the Carolinas, there appears to be no striking relationship between drought patterns and the topography or other readily identifiable meteorological characteristics of the Northeast.

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